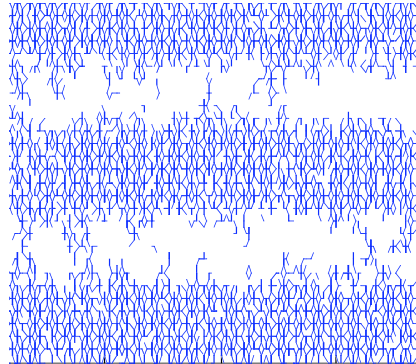




# The layering of polar firn and its linkage to air trapping

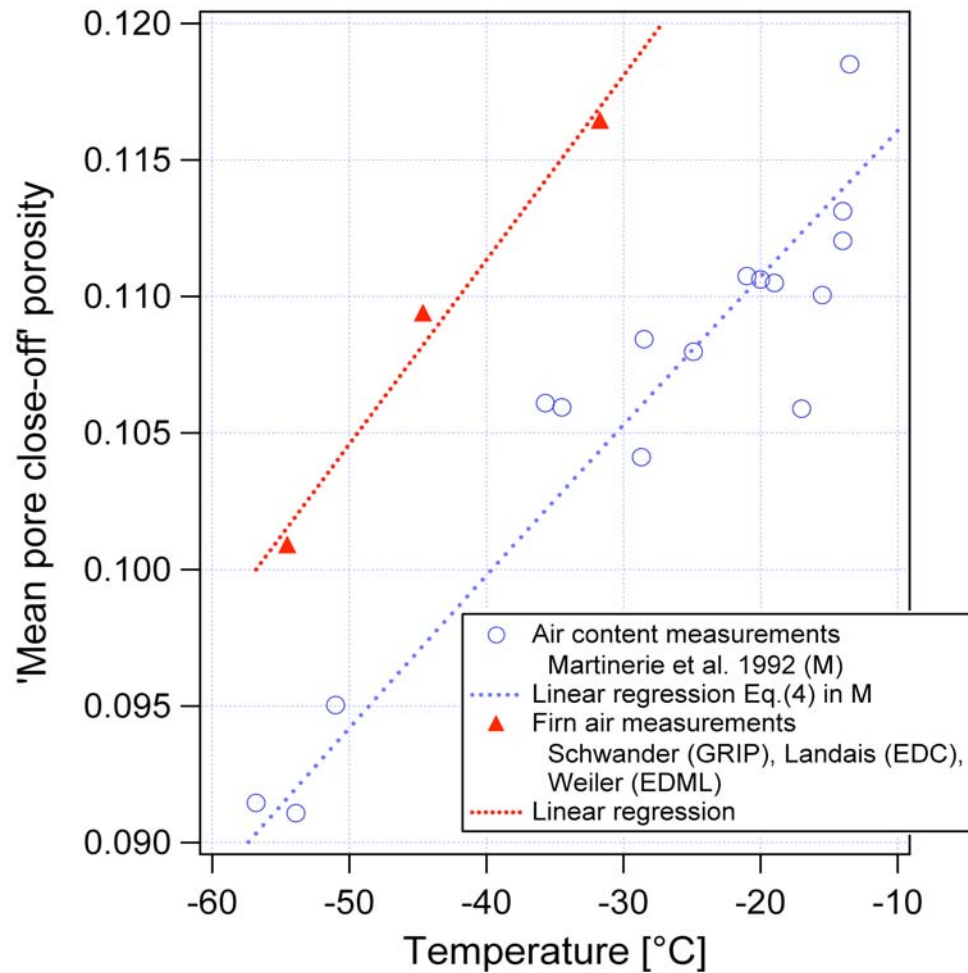


Johannes Freitag  
Sepp Kipfstuhl & Maria Hörhold

Alfred Wegener Institute (AWI)  
Foundation for Polar and Marine Research  
Bremerhaven



## The air content in polar ice



Linear relationship between pore close-off density (porosity) and temperature

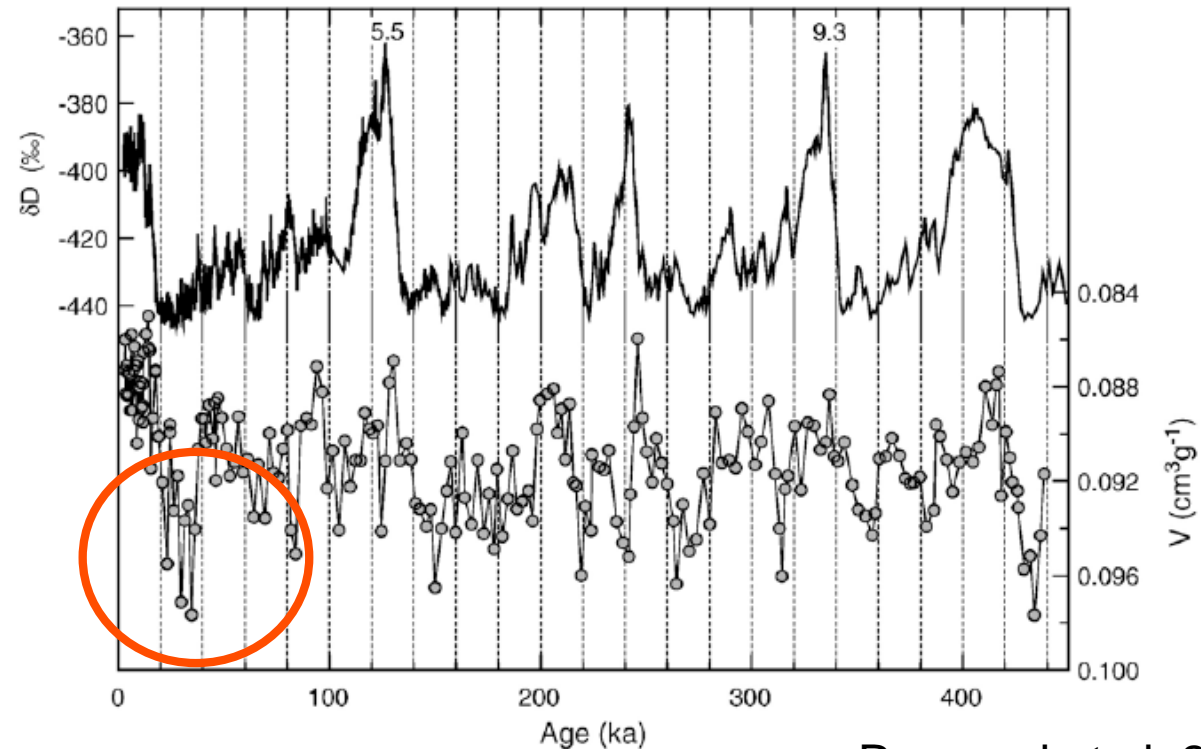
Martinerie et al. (1992)

Reason?

Relation to microstructure properties?



## Dome C (EDC) air content record



Raynaud et al. 2007

>80% of the variance in air content is not described by the  $n_{\text{crit}}(T)$ -relationship



## Methods: $\mu$ CT & Gamma-ray bench

### Results:

- **Direct observations of air trapping**

*An universal close-off porosity for polar firn?*

- **The evolution of layering with depth**

*Unexpected large layering at the firn-ice transition*

- **Simulations on stratified firn structures with a 3D-percolation model**

*Variations of mean close-off porosities*

## Summary



## X-ray-microfocus-computer-tomography ( $\mu$ CT) and Gamma-ray bench



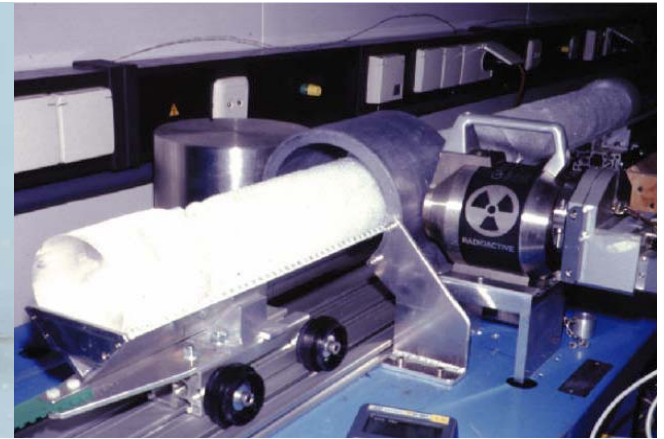
$\mu$ CT-SkyScan 1072,1074  
Resolution: 8-20 $\mu$ m, 40 $\mu$ m

Point measurements on the  
cm-scale

3D-volume processing:  
Topological parameters: Euler  
number, coordination number.



Field measurements  
Operation under cold  
conditions



Gamma-ray bench  
Vertical resolution:  $\sim$ 3mm

Continuous porosity  
measurements along  
the whole firn column  
(30 min per meter)



# Direct observation of air trapping



## Firn cores:

**B26**

North Greenland (77°06'N,49°06'E)

$T_{\text{mean}} = -30.6^{\circ}\text{C}$ , Accumulation rate= 179mm  
w.eq. a<sup>-1</sup>

**B36**

Kohnen station, EPICA, Dronning Maud Land  
Antarctica (75°00'S,00°04'E)

$T_{\text{mean}} = -44.6^{\circ}\text{C}$ , Accumulation rate= 62mm  
w.eq. a<sup>-1</sup>

**Dome C**

Dome C, EPICA, Antarctica (75°06'S,123°20'E)

$T_{\text{mean}} = -54.5^{\circ}\text{C}$ , Accumulation rate= 25mm  
w.eq. a<sup>-1</sup>



$$\text{Euler number } E := B - L + H \quad (1)$$

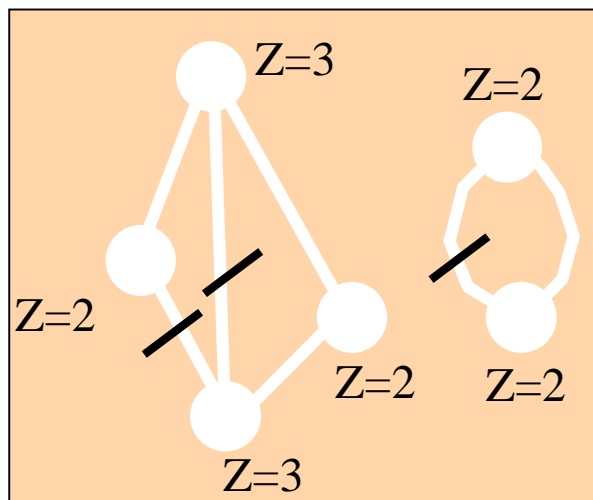
$$\text{Coordination number } \bar{Z} = \frac{2(B_p - E)}{B_p} \quad (2)$$

$B$ : Isolated objects

$L$ : Loops

$H$ : Holes

$B_p$ : Potential objects



$$\begin{aligned} B &= 2 \\ L &= 3 \\ H &= 0 \end{aligned}$$

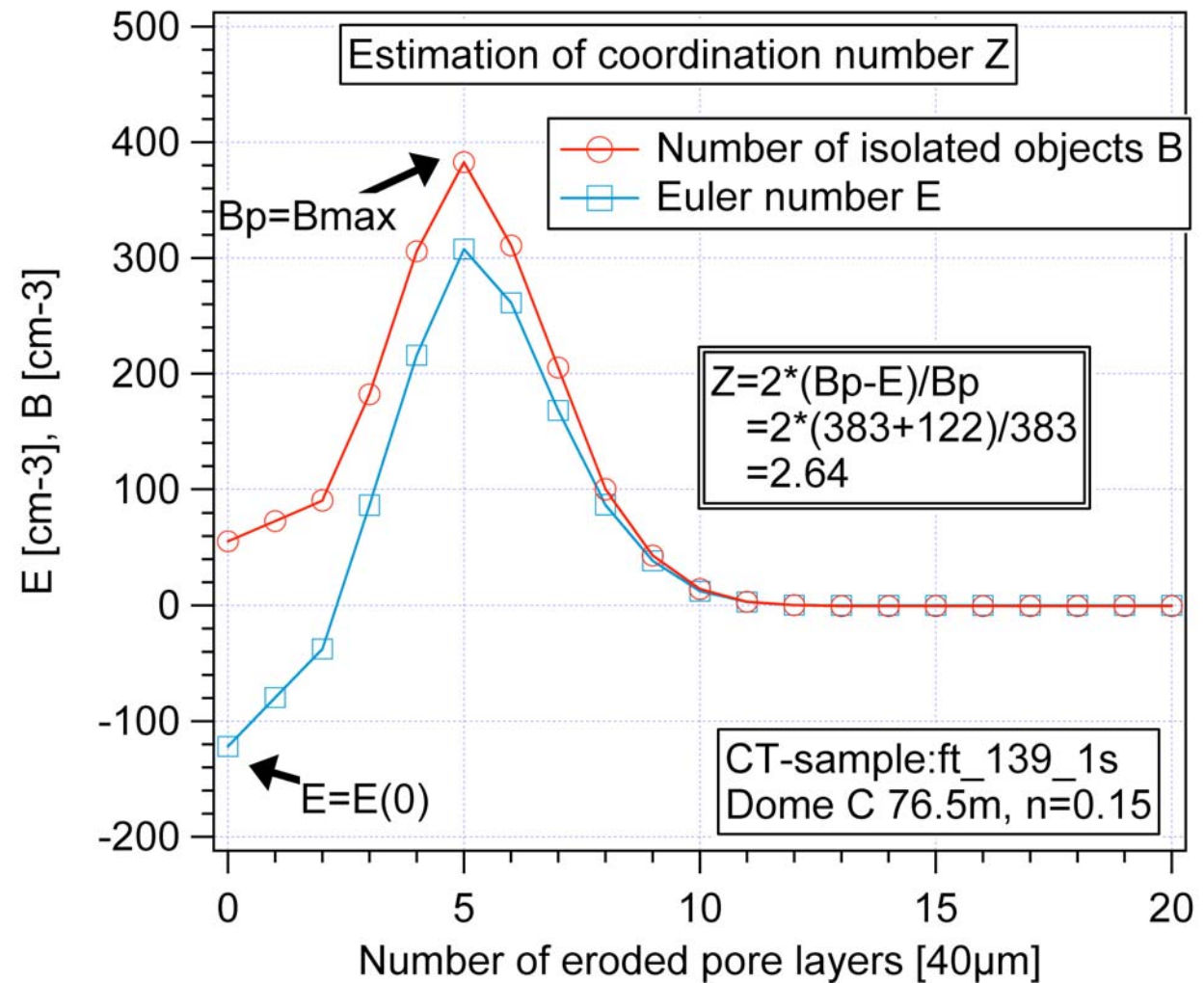
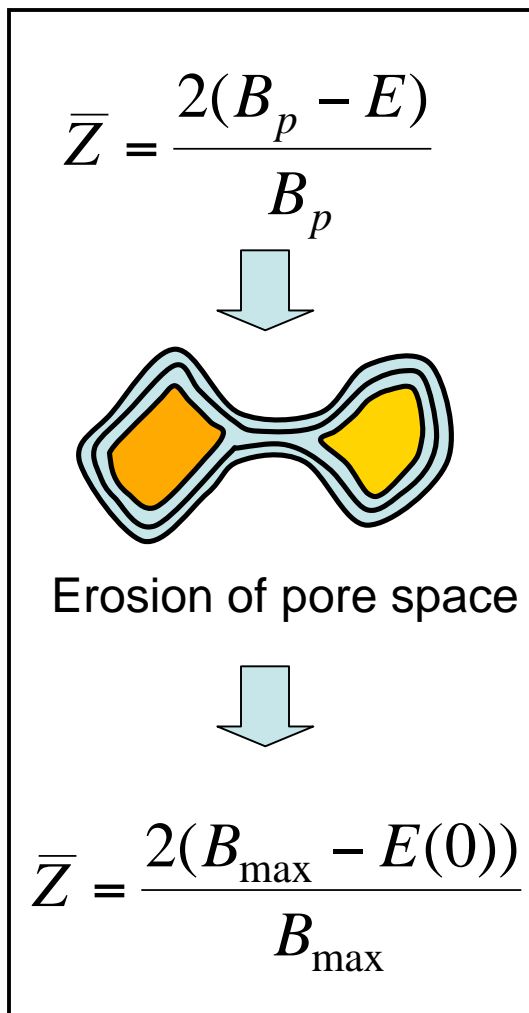
$$\begin{aligned} (1) \quad &\Rightarrow E = B - L = 2 - 3 = -1 \\ &\bar{Z} = \frac{(2 + 2 + 2 + 2 + 3 + 3)}{6} = 2.33 \end{aligned}$$

$$B_p = 6$$

$$\begin{aligned} (1,2) \quad &\Rightarrow \bar{Z} = \frac{2(6 - (-1))}{6} = 2.33 \end{aligned}$$



## 3D-image processing

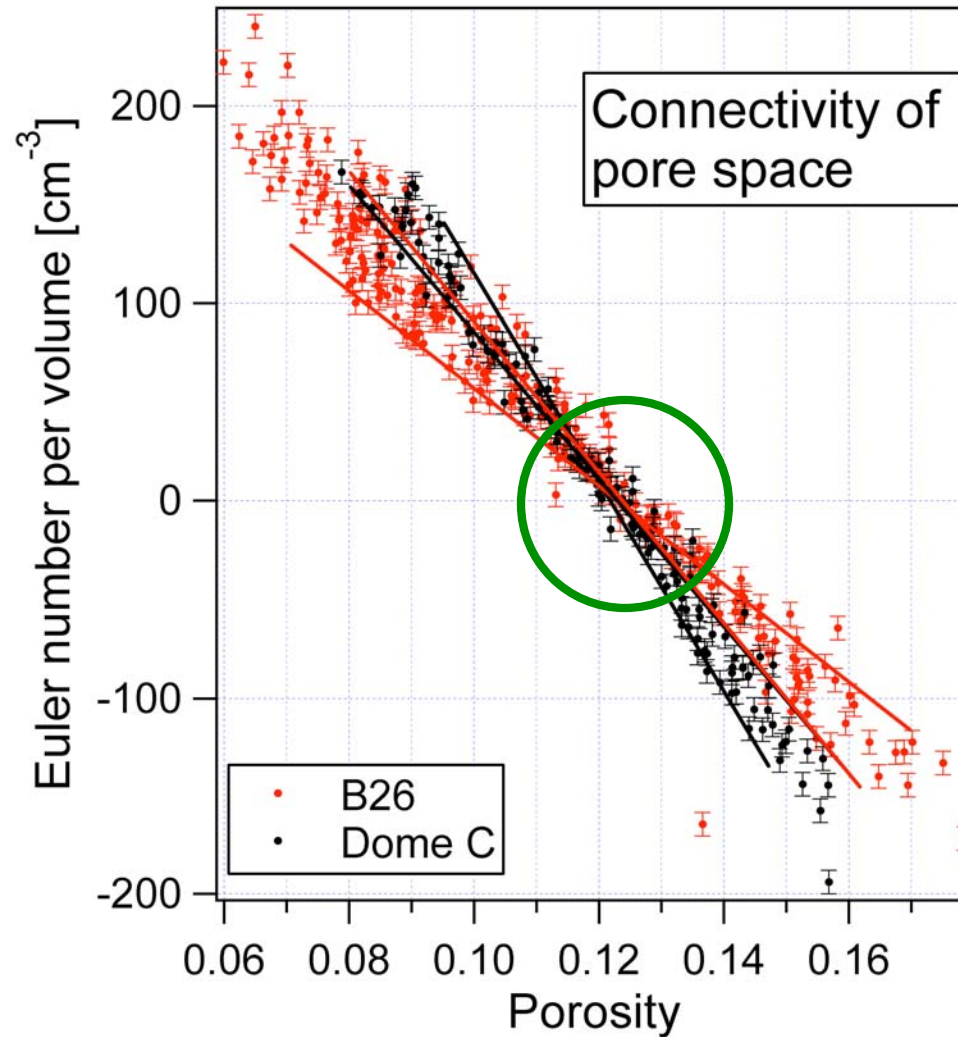


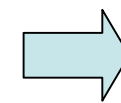
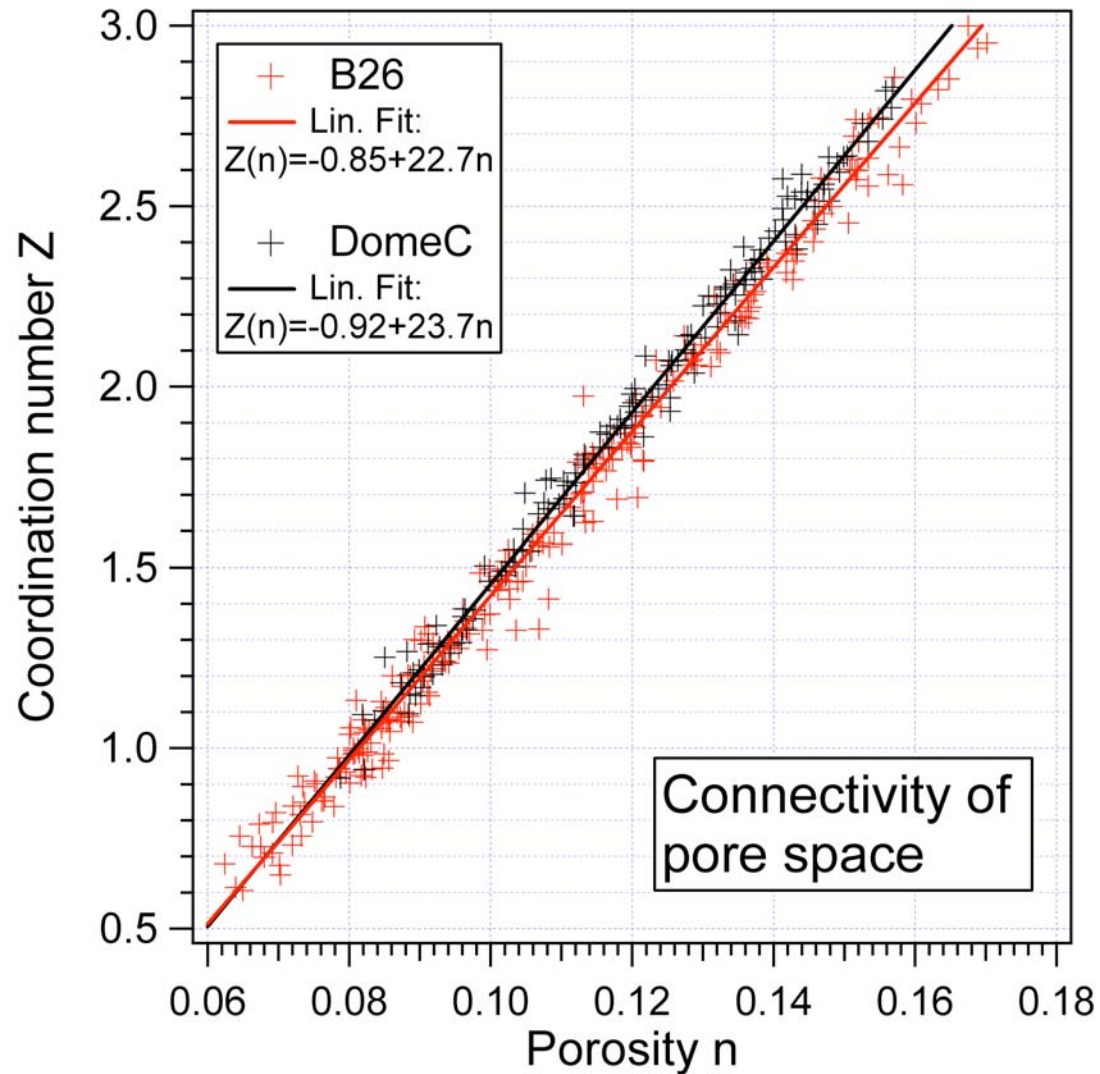


$E \sim B$  ( $\sim$  bubble density)

$$E=0 \xrightarrow{(2)} Z=2$$

$E \sim L$  ( $\sim$  Loops)

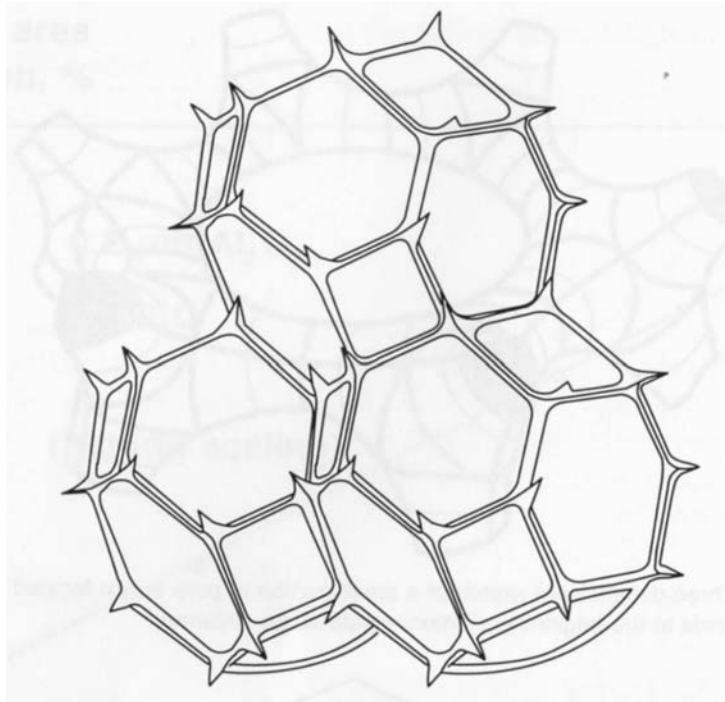




$Z(n)$   
linear relationship



## Pore close-off as a percolation problem



Fully occupied lattice ( $p=1$ ):

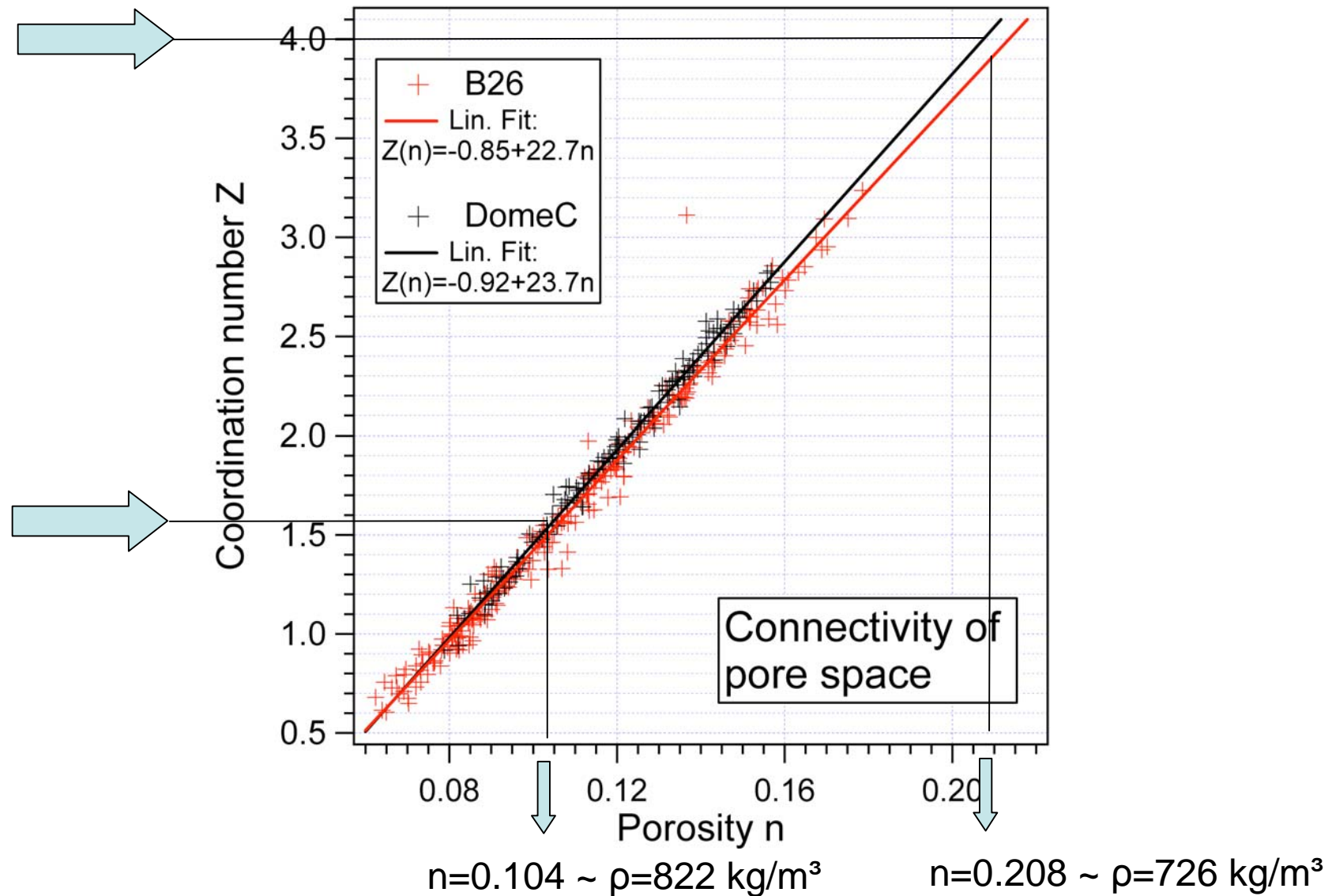
$$Z=4$$

Lattice at the threshold from permeability to impermeability (percolation threshold  $p_{\text{crit}} \sim 0.4$ )

$$Z \approx 1.6$$

Best lattice geometry for firn:  
Sintered spheres (tetrakaidecahedrons),  
grains are placed on a bcc-lattice

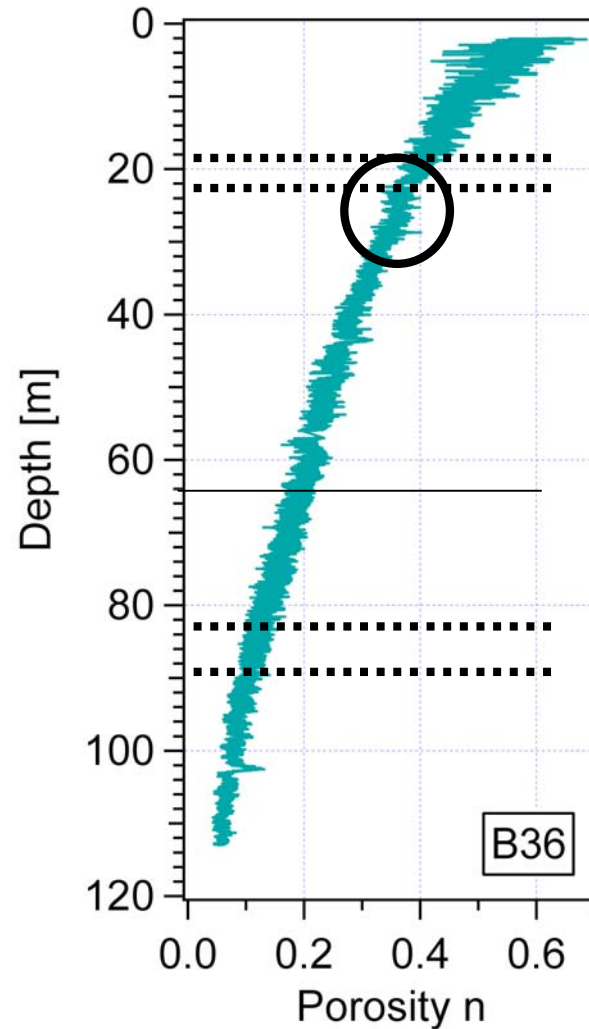
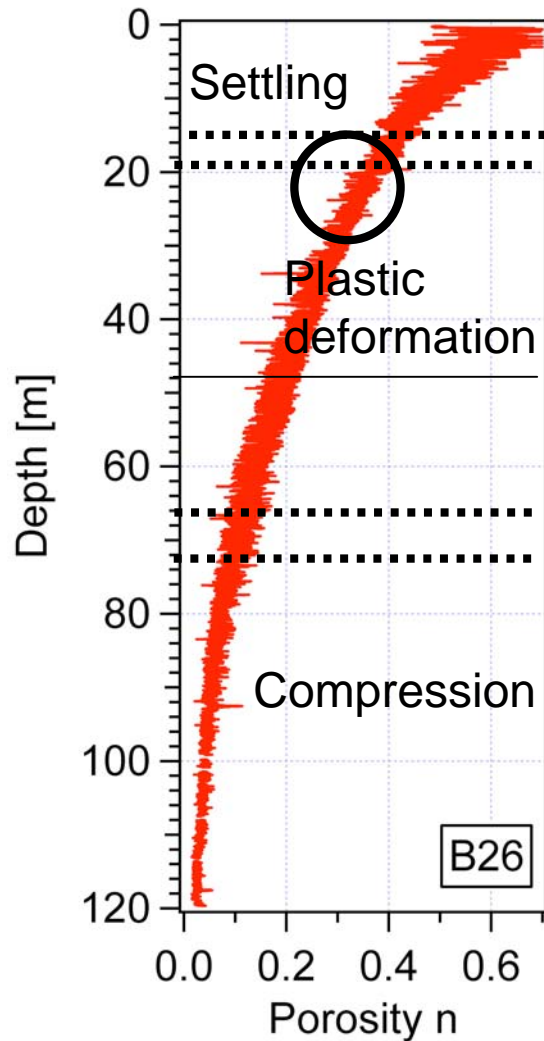
$$Z(p) \approx \text{linear}$$





# The evolution of layering

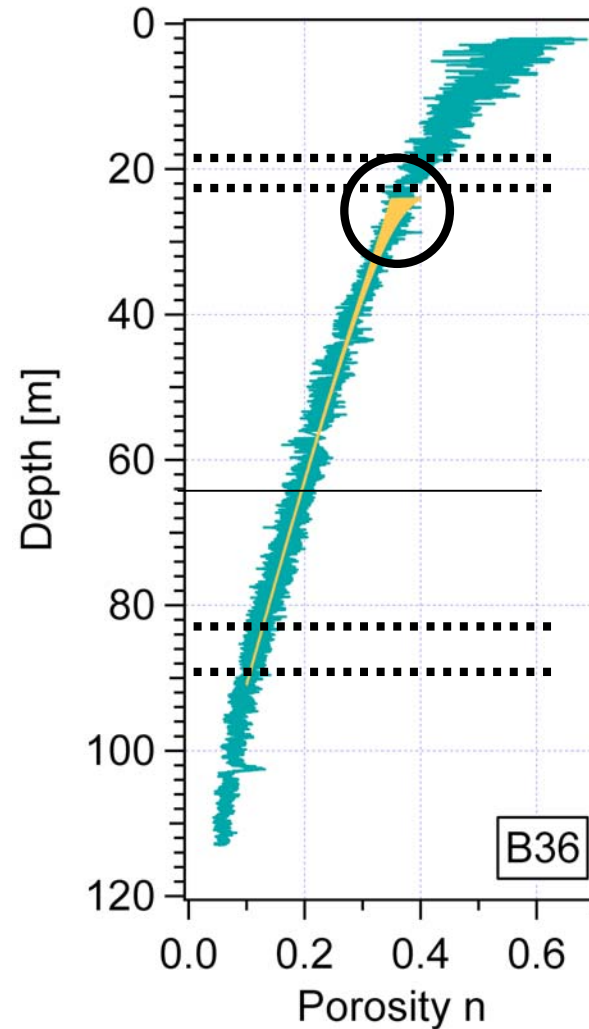
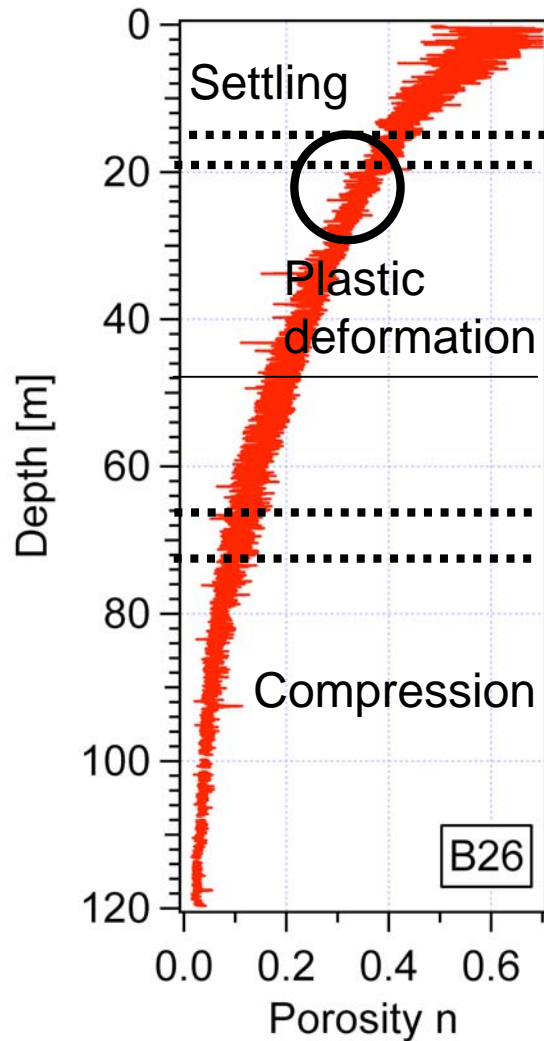
Image: S. Kipfstuhl

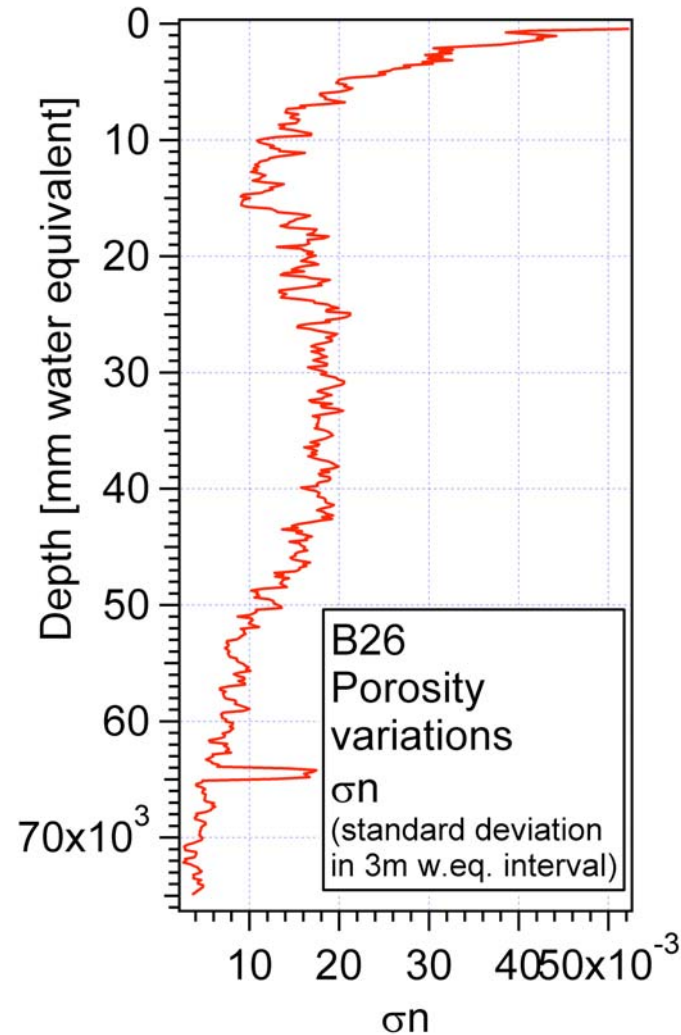
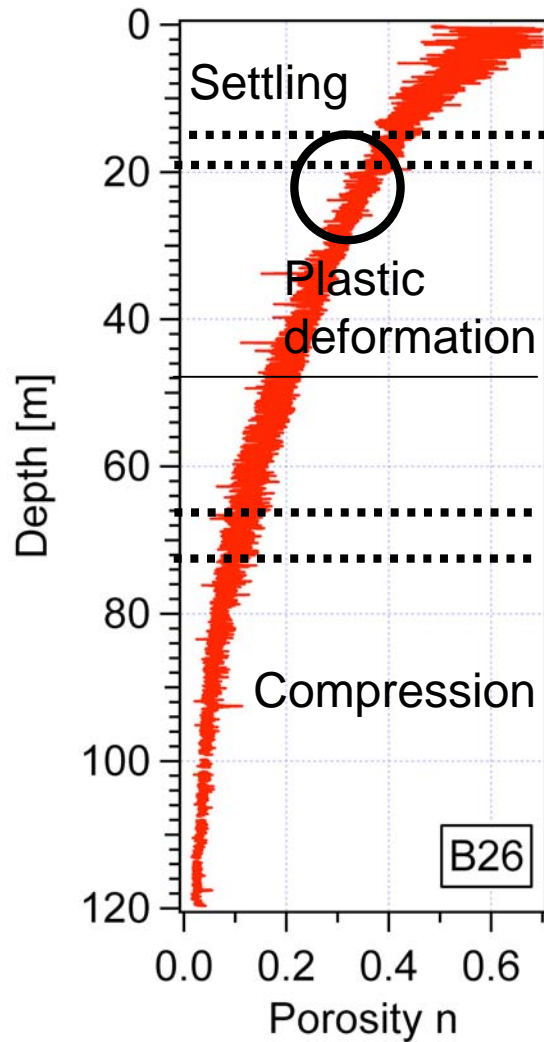


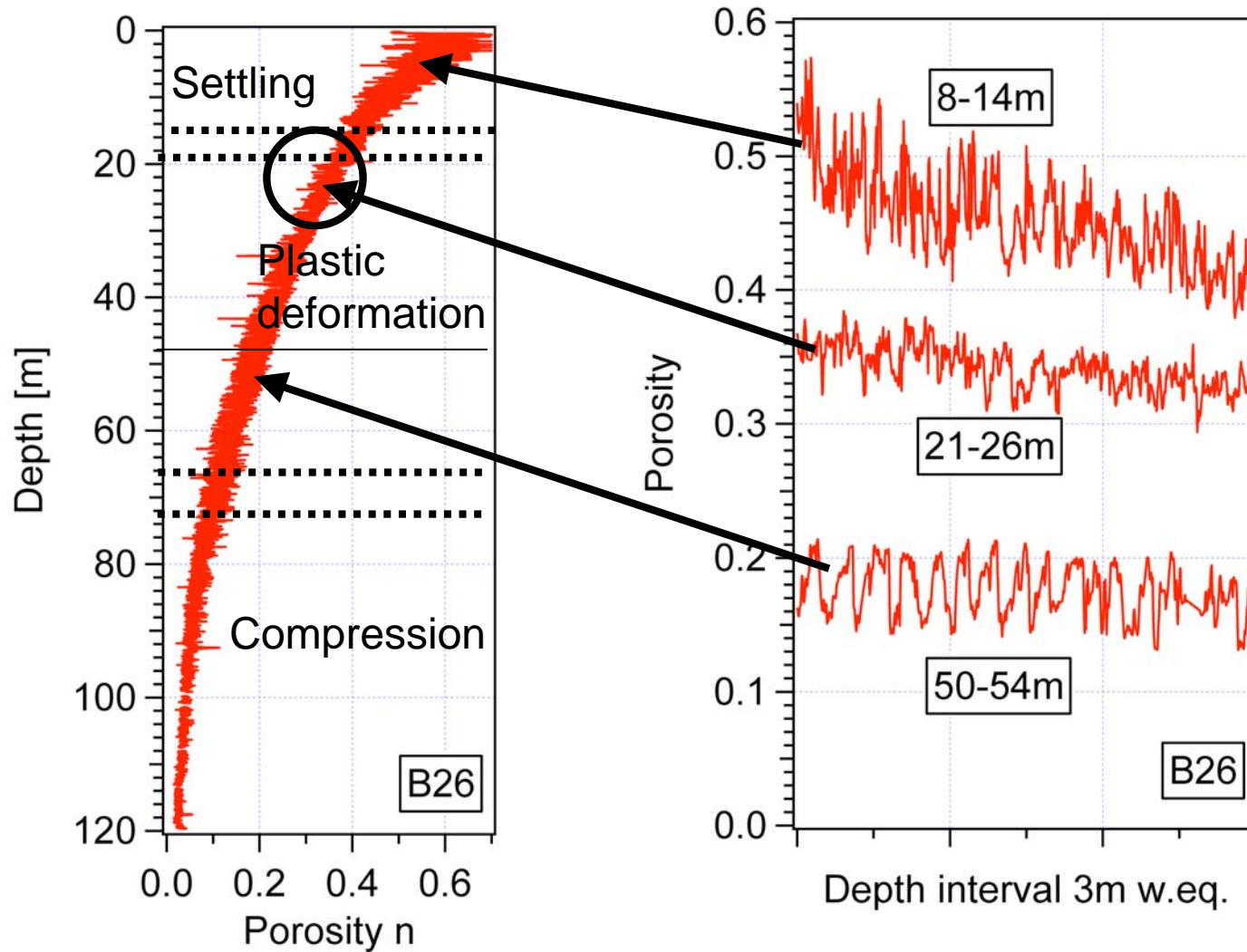
550 kg/m<sup>3</sup>      n=0.4

730 kg/m<sup>3</sup>      n=0.2

Firn-ice transition  
820 kg/m<sup>3</sup>      n=0.1

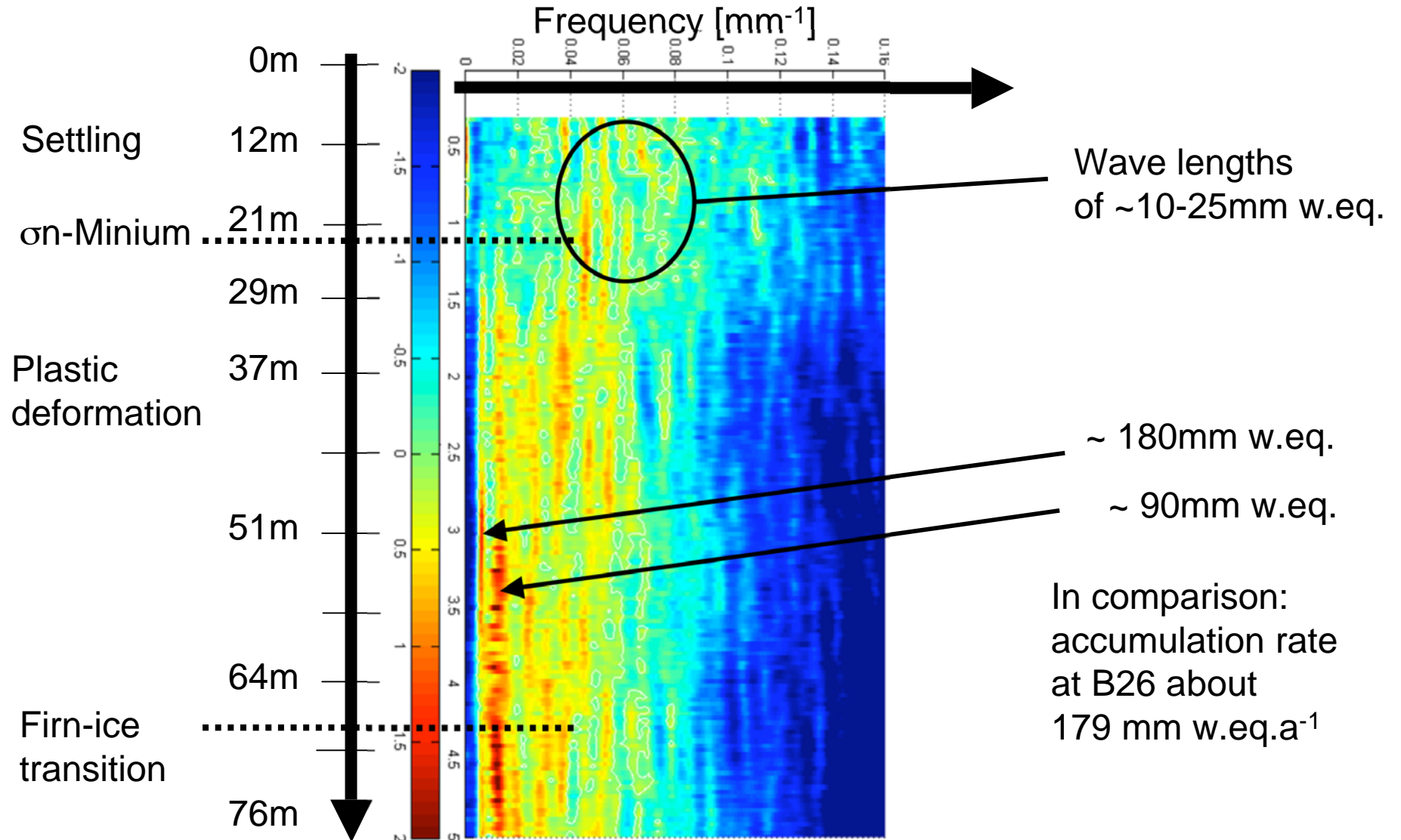


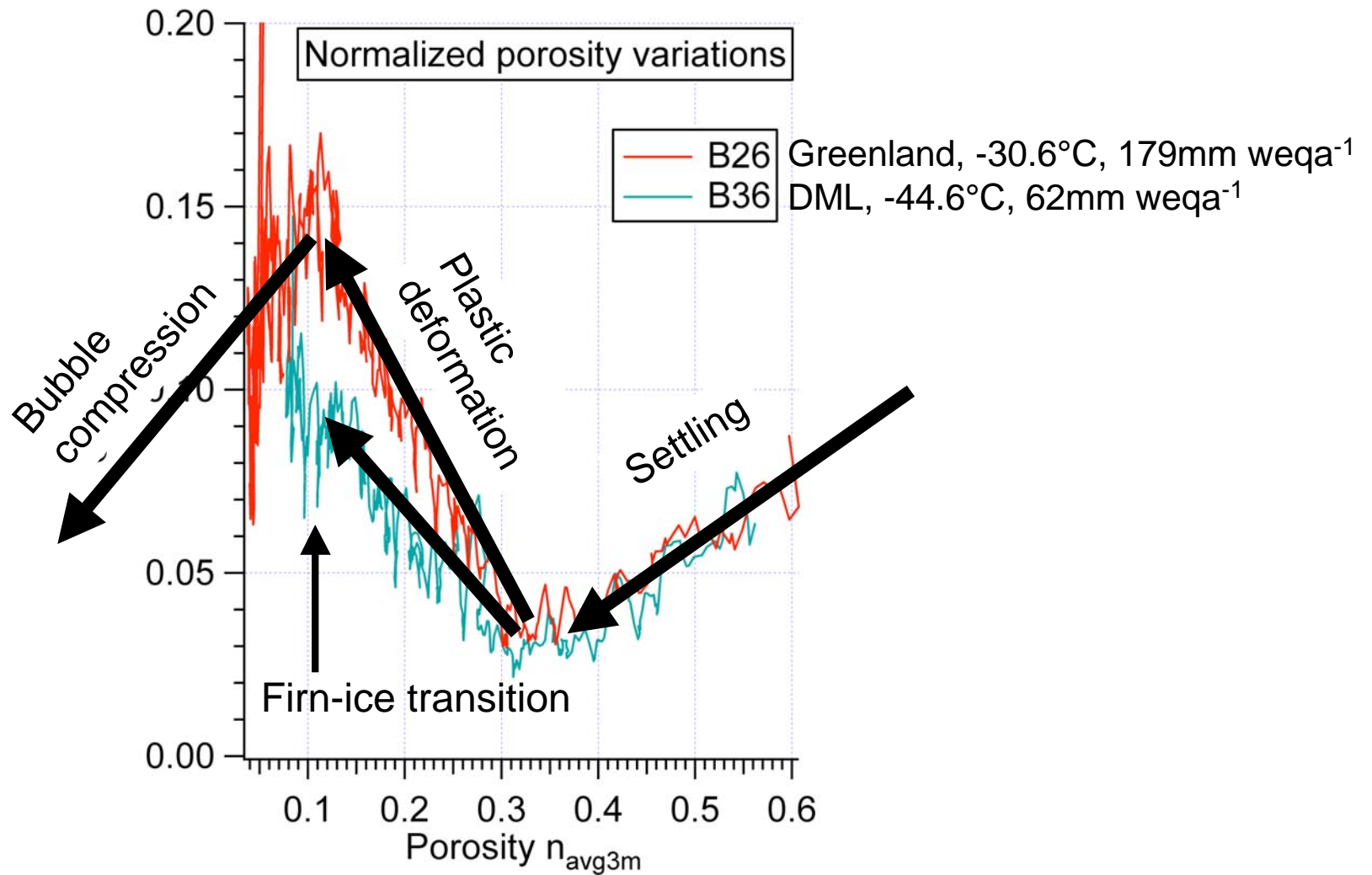






Evolutionary power spectrum of porosity (B26)







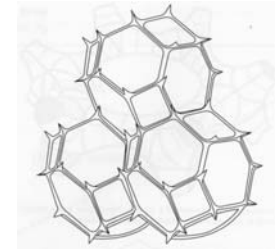
# Percolation on stratified firm



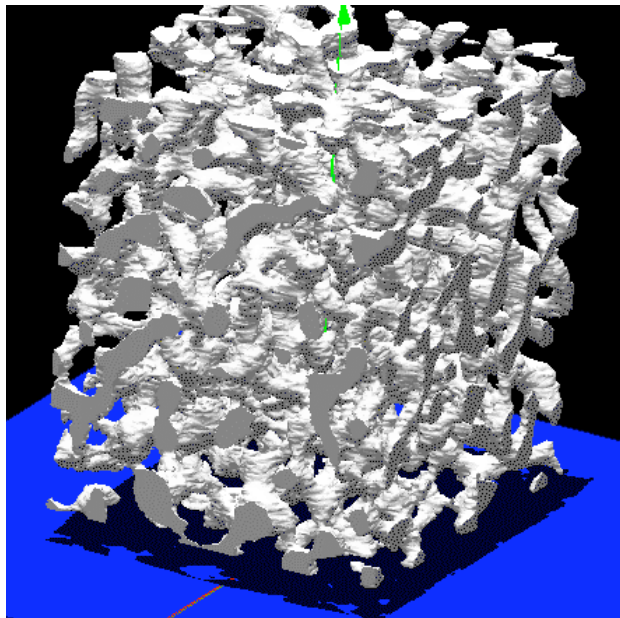
## Adapting the percolation model on real firn pore structures

Step 1

Introduction of a length scale  
via the Euler number in fully occupied structures ( $Z=4$ ):

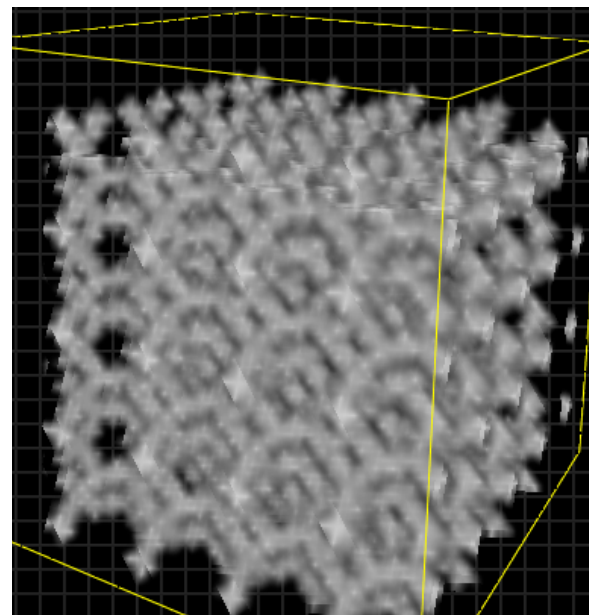


Pore structure at  $n \approx 0.2$



B26, 40m,  $E \approx -650$ ,  $12 \times 12 \times 12 \text{ mm}^3$

Lattice structure at  $p=1$



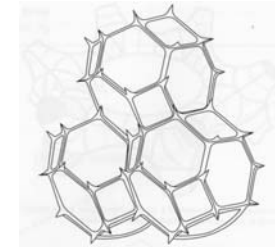
BCC-lattice,  $4 \times 4 \times 4$  units,  $E \approx -680$

$E(\text{firn}) \approx E(\text{lattice}) \implies 1 \text{ bcc unit} = 3 \text{ mm (B26)},$   
 $= 2.7 \text{ mm (DomeC)}$



## Adapting the percolation model on real firn pore structures

Step 2

Introduction of layering  
via measured porosity variations at the firn-ice transitions

Assumption

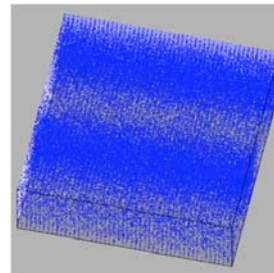
Occupation probability varies periodically  
in vertical direction (laterally homogenous)Amplitude:  $\sigma n$ 

Wave lengths: year's accumulation

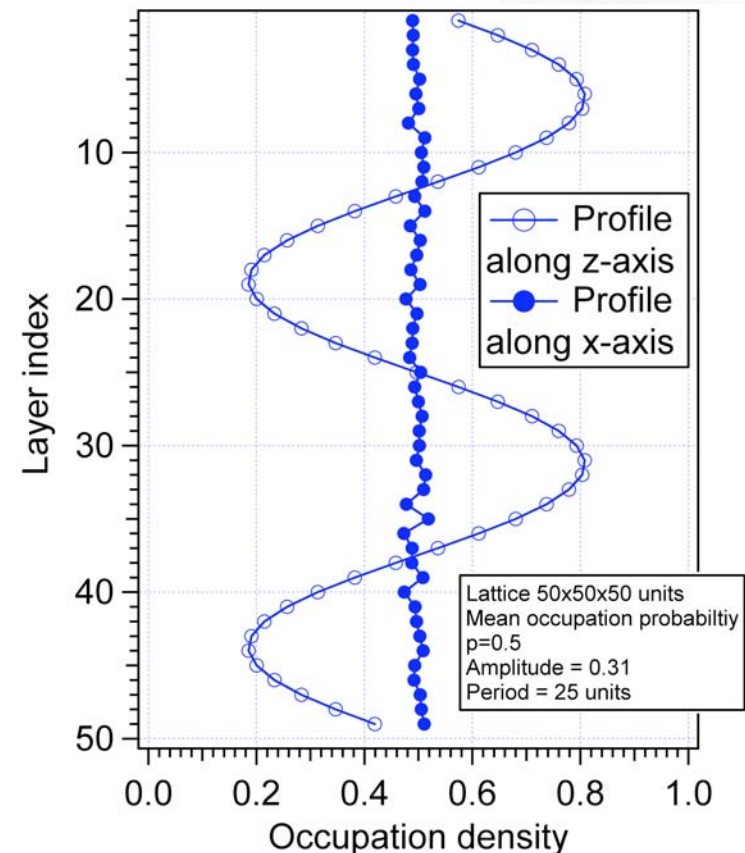
Conversion:

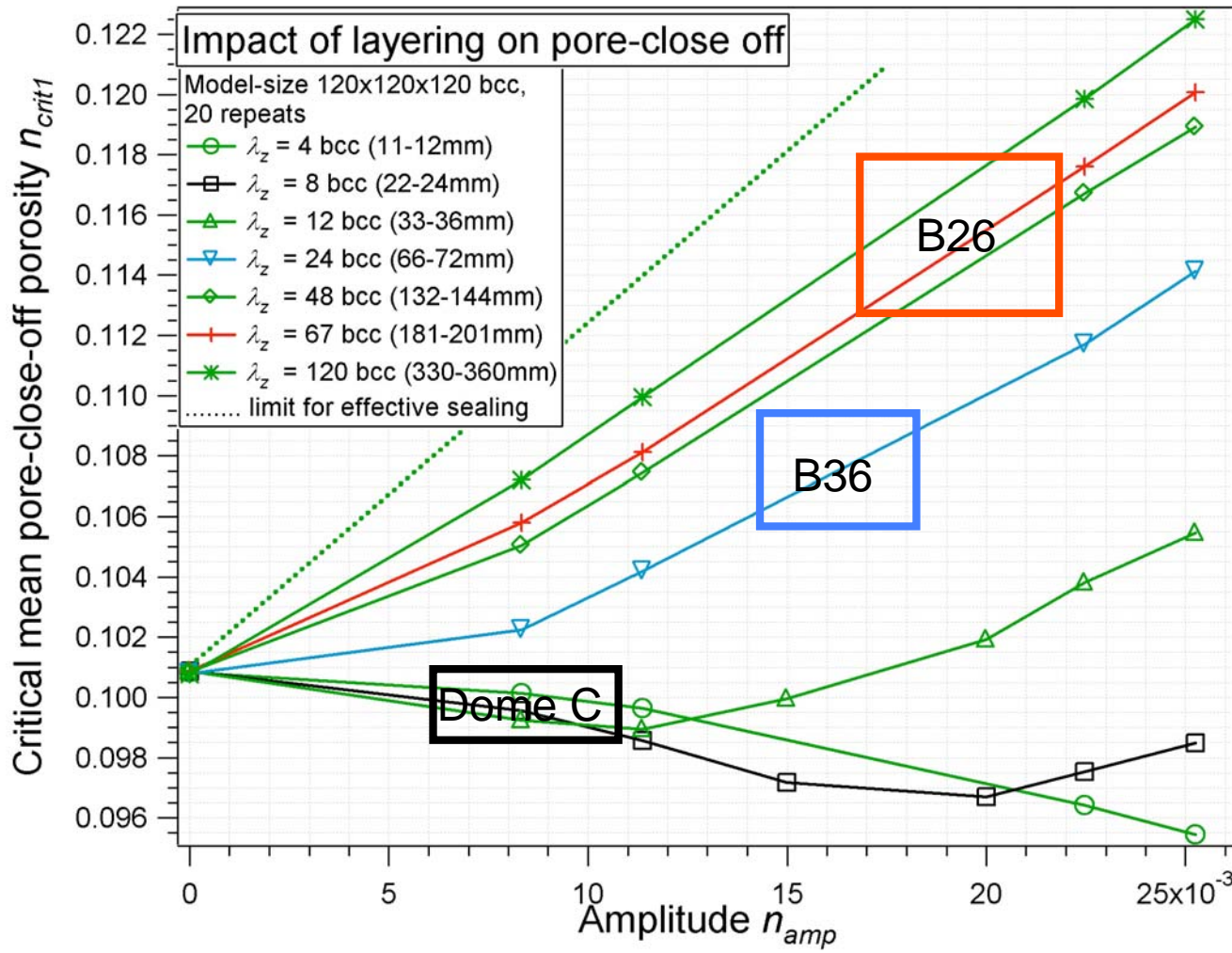
$$p(n) = -0.22 + 5.8n$$

(from the  $n(Z)$  and  $p(Z)$ -relation)



$$p(x, y, z) = p_{avg} + p_{amp} \sin \frac{2\pi z}{\lambda_z}$$





Firn air Pumping measurements

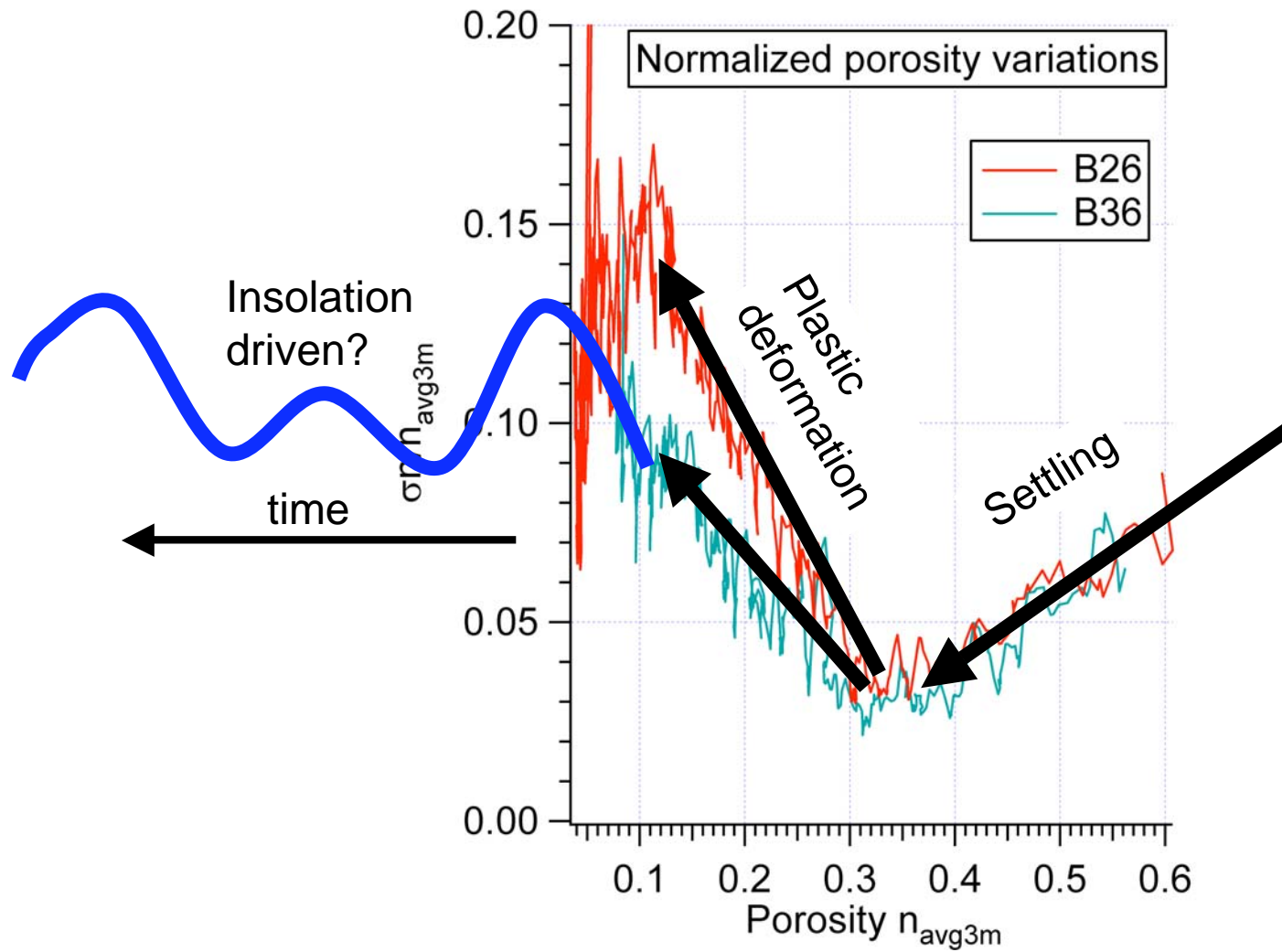
GRIP

B36

Dome C



- Evidence for an universal close-off porosity for polar firn
- The 3D-percolation approach explains two critical porosities/densities ( $0.101 / 820 \text{ kg/m}^3$  and  $0.21 / 730 \text{ kg/m}^3$ ) of polar firn
- Unexpected large layering at the firn-ice transition
- Layering at firn-ice transition mainly results from different densification rates in seasonal layers in the stage of plastic deformation (stage II). The variations start to synchronize with the seasonal cycle. Are they chemically induced?
- Recent variations of mean close-off porosities at different sites can be exclusively explained by layering (amplitude and wave length are important)



Thank you



J. Freitag



Firn Workshop, Dartmouth College, March 10-11, 2008







## Open questions concerning the firn-ice transition

The physical link between insolation and N<sub>2</sub>/O<sub>2</sub>, air content (Duval?, Kawamura,..)

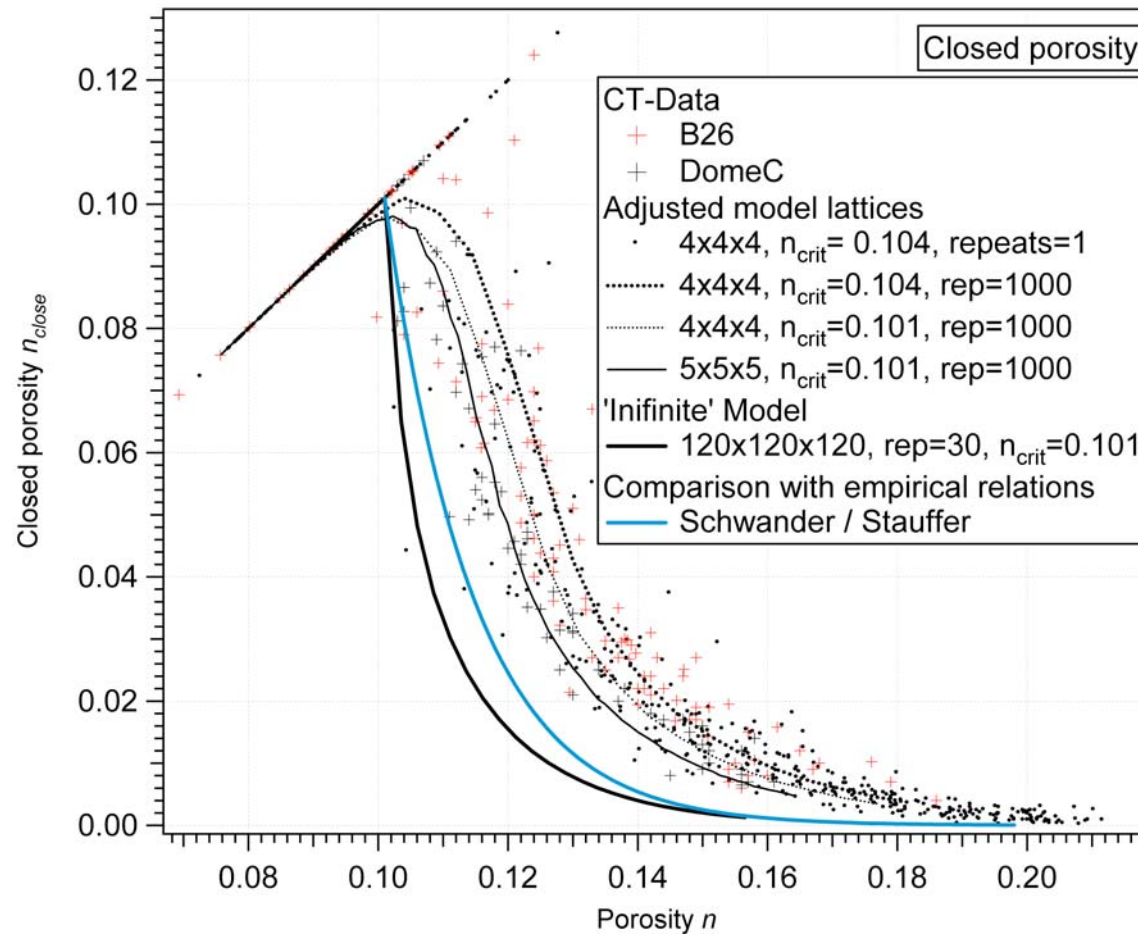
Large discrepancy between modelled close-off depths and  $\delta^{15}\text{N}$  inferred lock-in depths under Glacial conditions in low accumulation sites (Landais et al, 2006).

The nature of the non-diffusive zone

Why is the empirical relationship between pore close-off density (porosity) and temperature found by Martinerie et al. (1992) only valid for recent firn but not for paleo firn?

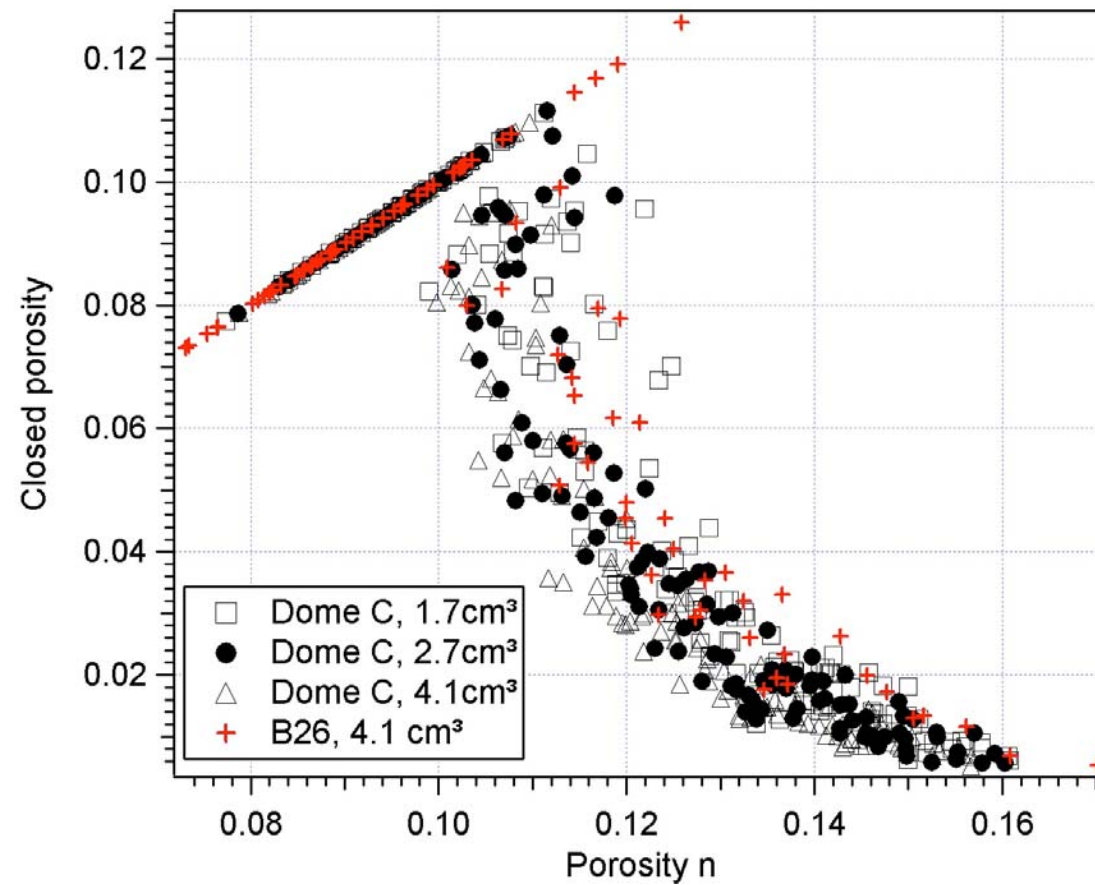


## Closed porosity: model versus data





## Closed porosity: size effect of data



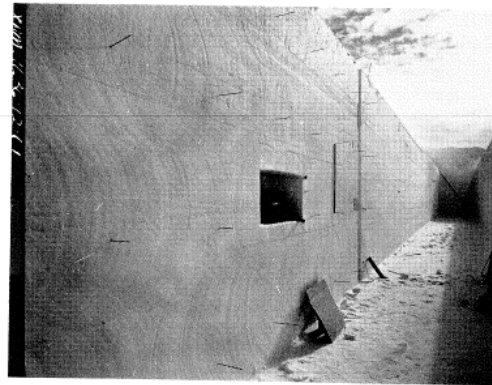


Fig. 6b. The first 6 meters of the study area with the layering accentuated by sunlight.

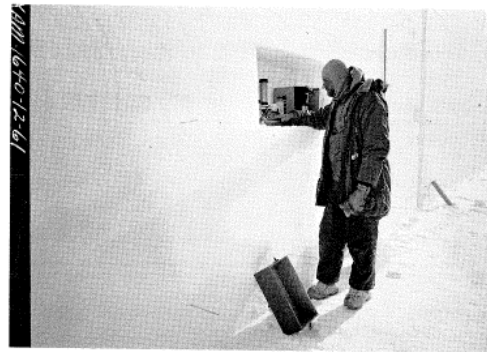


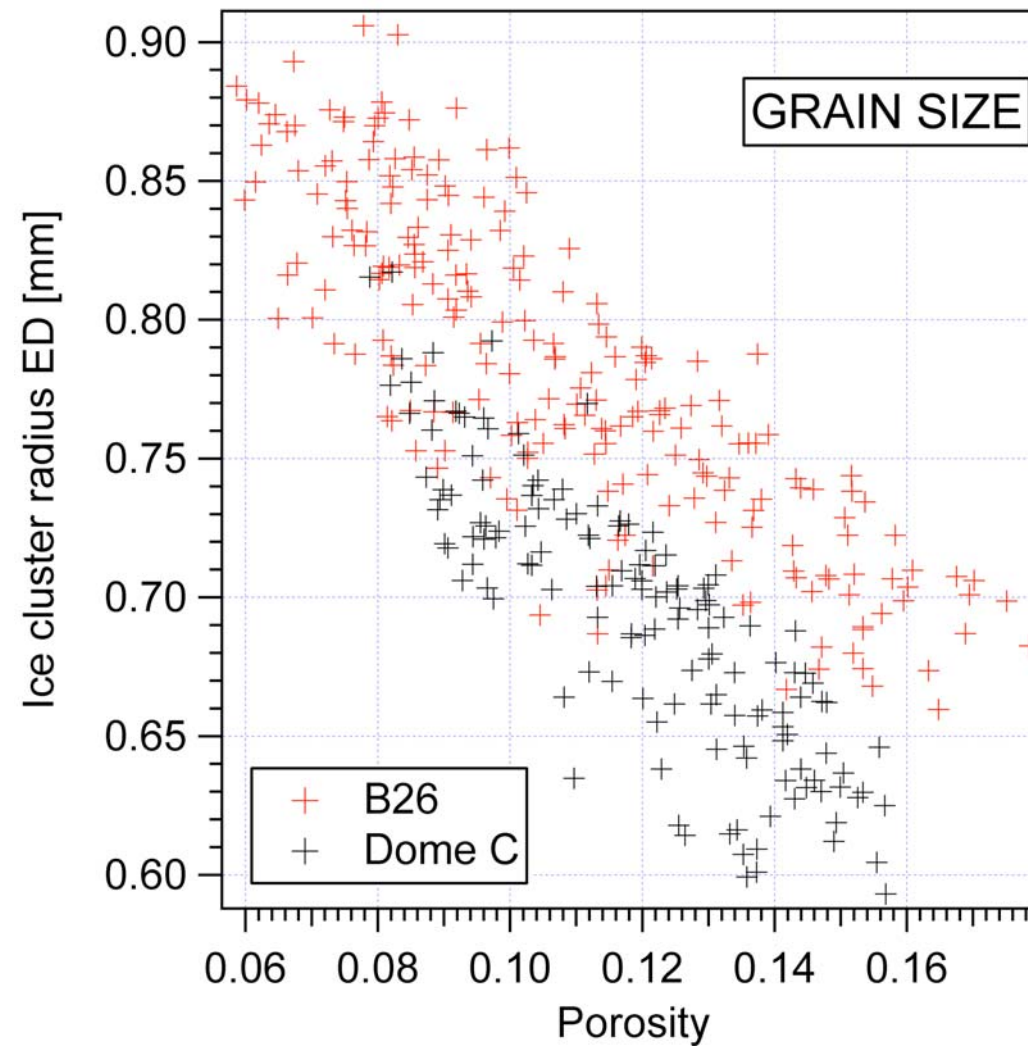
Fig. 6d. Density samples cut from the pit wall were weighed on a triple-beam balance mounted in the shelf shown here.

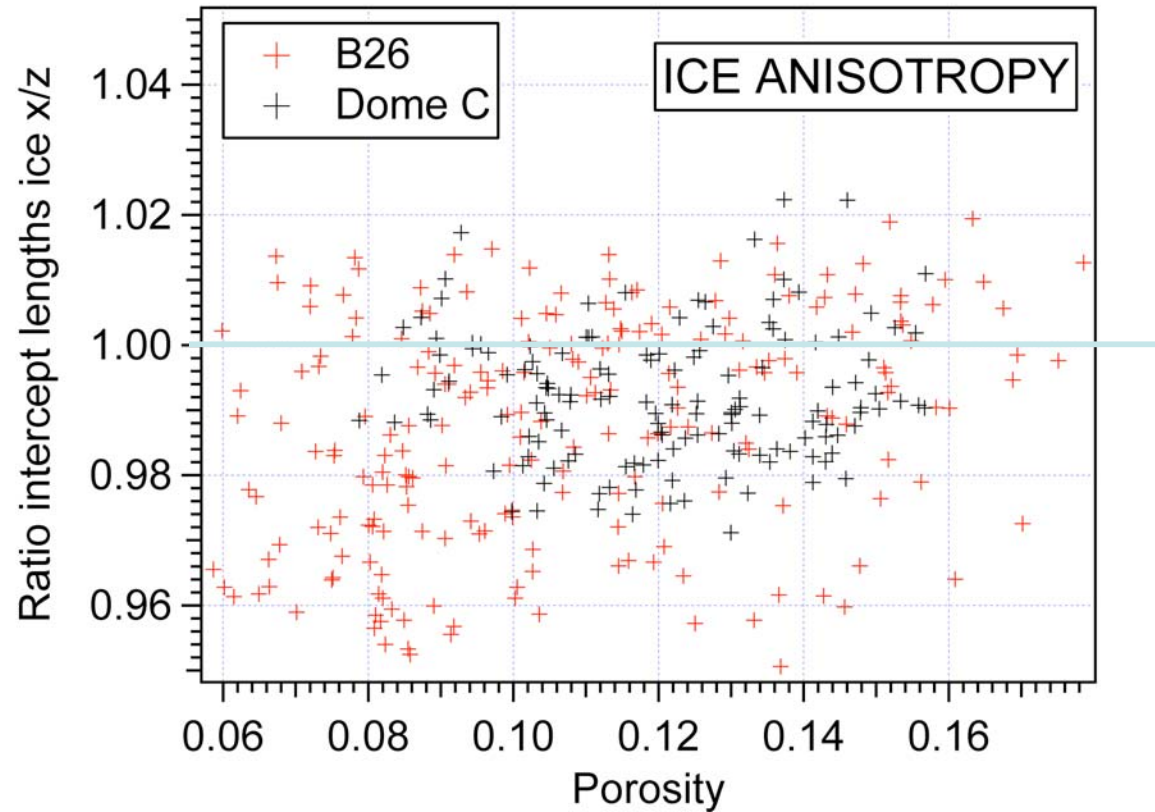
... 1961

Stratigraphic studies at  
Byrd Station, Antarctica,  
by C. Benson.

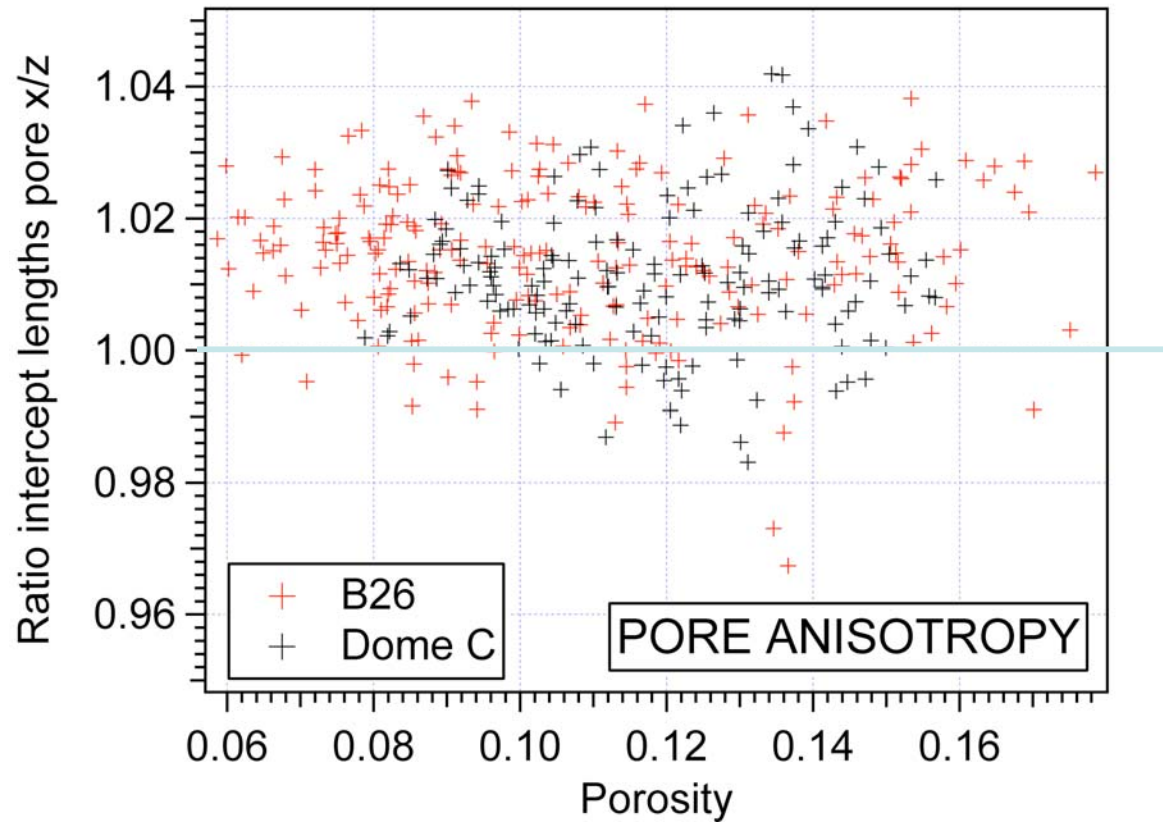
... 2008

Why are we still  
interested  
on firn layering?

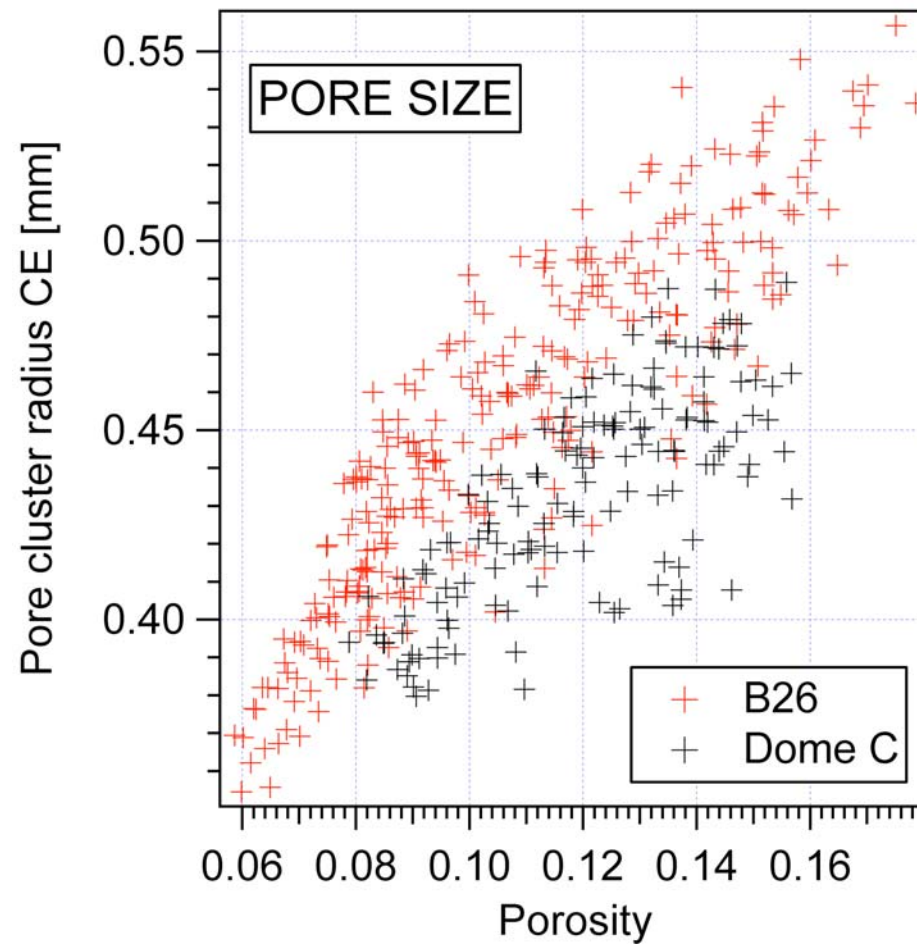


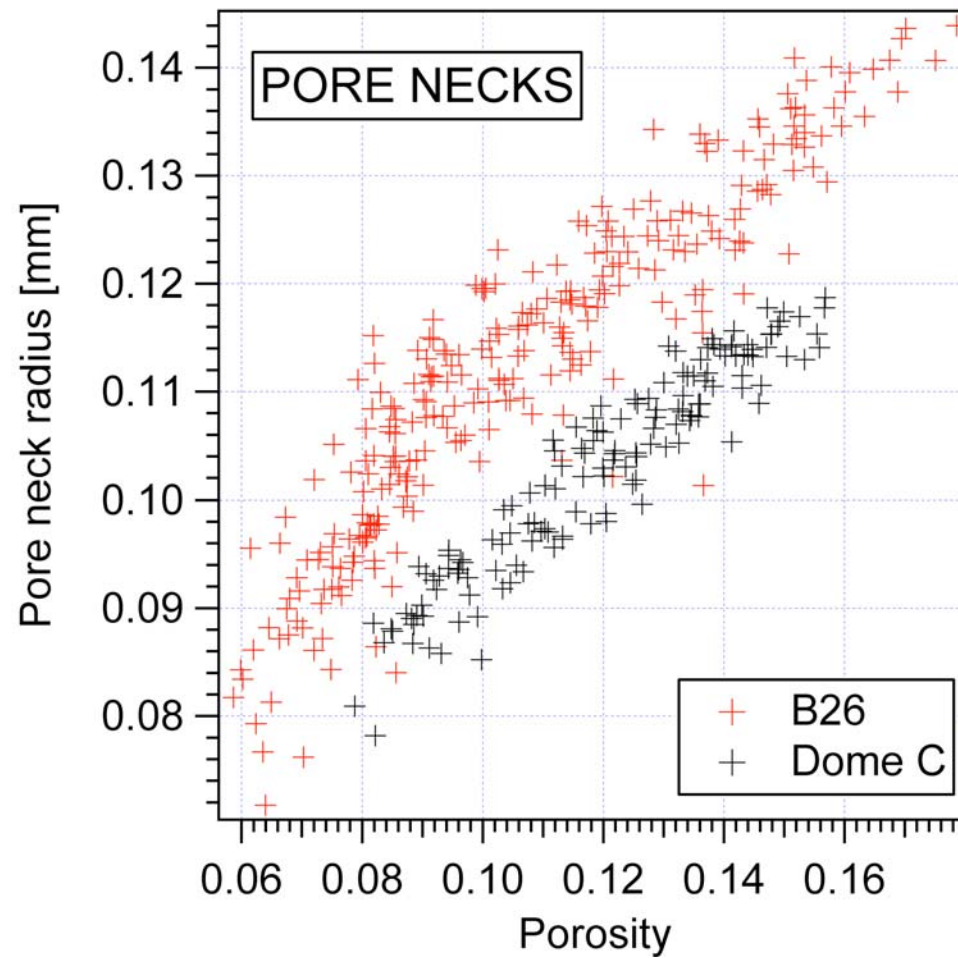


Ice phase is vertically elongated



Pore phase is vertically compressed







## X-ray-microfocus-computer-tomography ( $\mu$ CT)



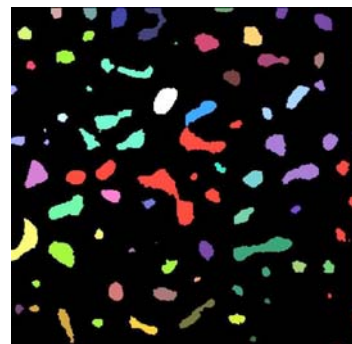
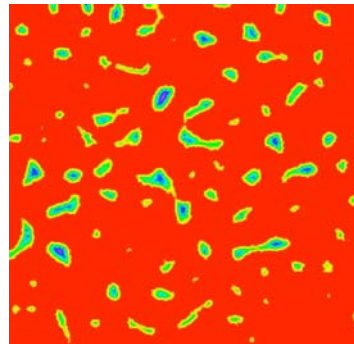
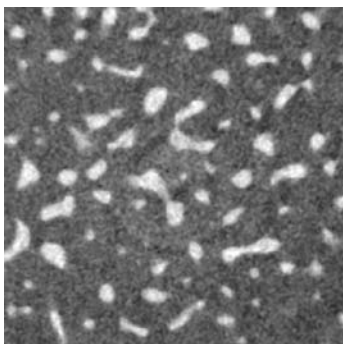
SkyScan 1072  
Resolution 8-20 $\mu$ m



Field measurements



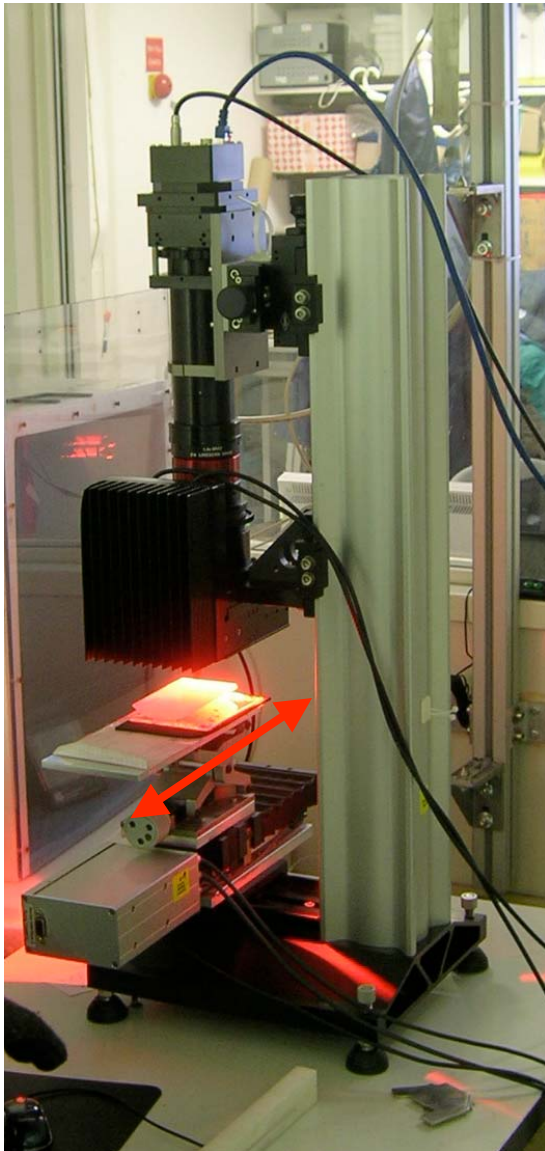
SkyScan 1074  
Resolution: 40 $\mu$ m



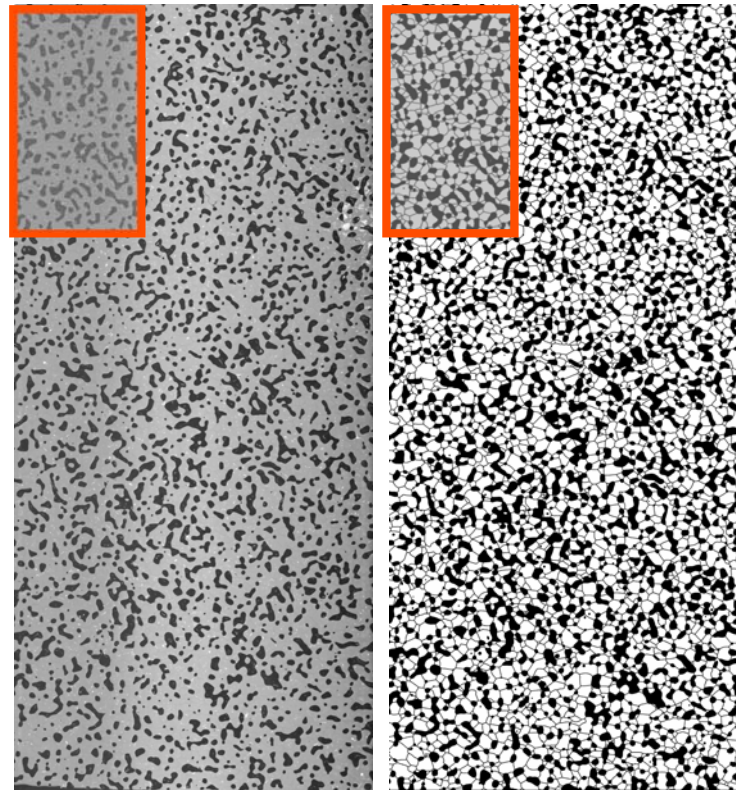
3D-volume processing:  
open and closed porosity,  
pore and ice cluster size  
distributions, neck radii,  
coordination number,  
anisotropy, specific surface  
density, mean curvature...



# Grain-size-Scanner

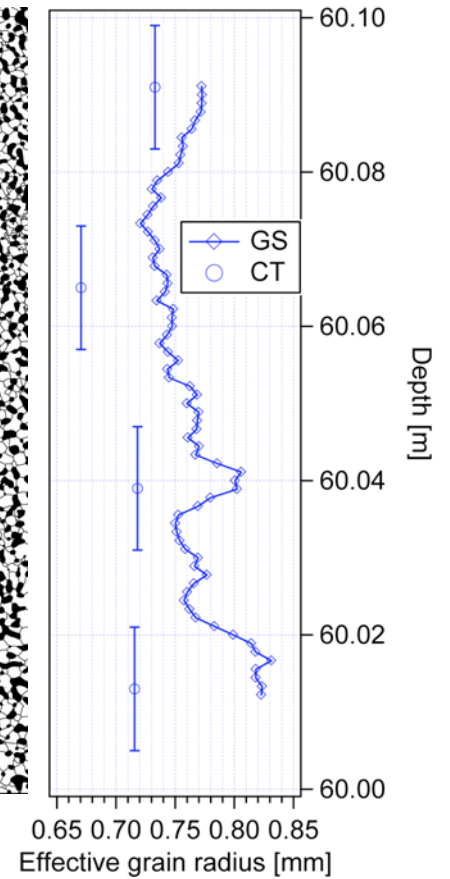


10 cm



EDML, 61m

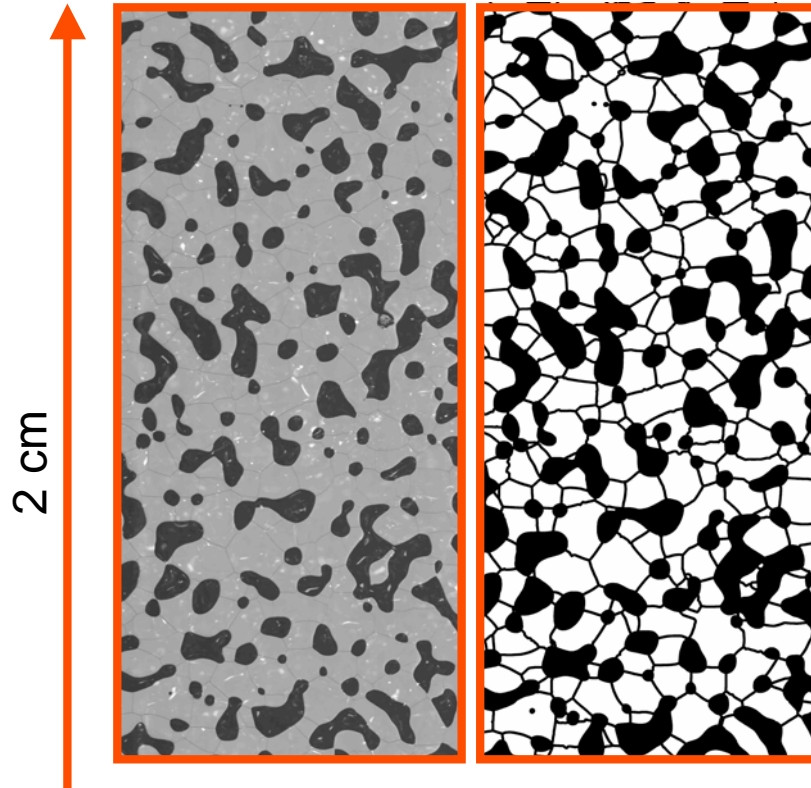
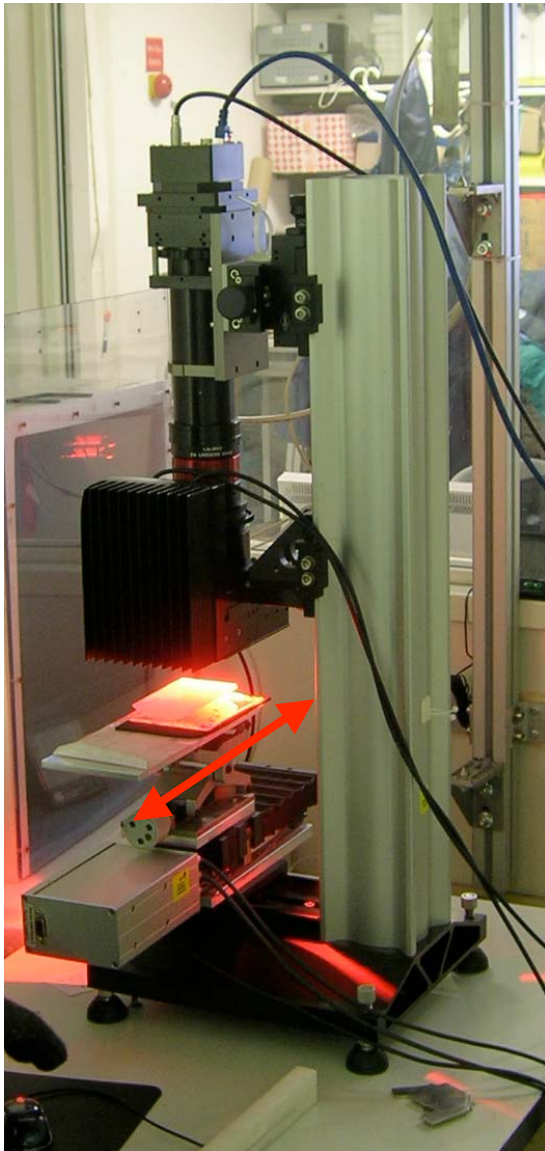
## Intercomparison



Resolution 5 $\mu$ m, 4cm scan width, v=1cm/s  
(Schäfter&Kirchhoff, Hamburg)

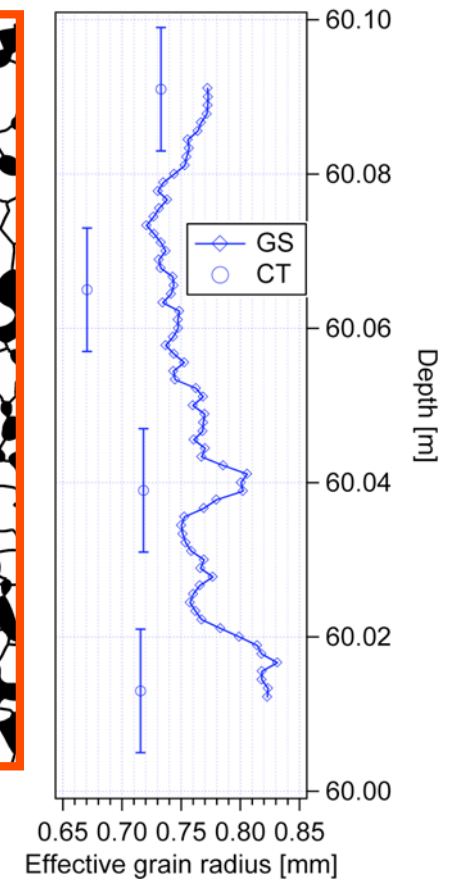


# Grain-size-Scanner



EDML, 61m

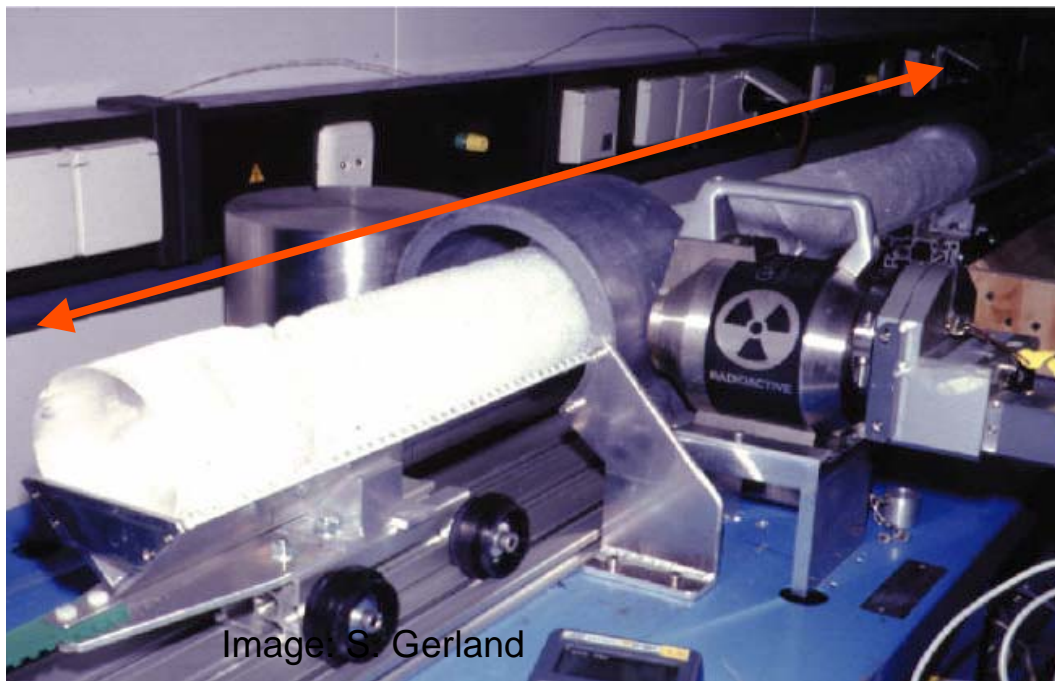
## Intercomparison



Resolution 5 $\mu$ m, 4cm scan width, v=1cm/s  
(Schäfter&Kirchhoff, Hamburg)



## Gamma-bench for porosity measurements (GAMB)



Vertical resolution: ~3mm

### Intercomparison

